

# Approximating the complex-valued Green-tensor amplitude by a real-valued Green-tensor amplitude

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## Summary

The paper is devoted to the approximation of the complex-valued Green-tensor amplitude by a real-valued Green-tensor amplitude with a phase shift. This approximation is required in the method of Jan Šílený for determining the real-valued seismic moment tensor from the maximum real-valued vectorial amplitude picked in the polarization diagram.

## Keywords

Wave propagation, elastic waves, Green tensor, amplitude, seismic moment tensor.

## 1. Introduction

The tensorial amplitude of the S-wave Green tensor is often complex-valued, especially if the wavefield is recorded at the Earth surface (Červený, 2001).

The method of Šílený & Milev (2008), Šílený et al. (2009) and Horálek & Šílený (2013) for determining the real-valued seismic moment tensor from the maximum real-valued vectorial amplitude picked in the polarization diagram requires the approximation of the complex-valued Green-tensor amplitude by a real-valued Green-tensor amplitude with a phase shift. When studying this approximation, we shall assume the ray-theory approximation of the Green tensor expressed in terms of the  $3 \times 3$  tensorial amplitude and travel time. The approximation is applicable to Green-tensor amplitudes in both isotropic and anisotropic heterogeneous media, including the isotropic ray theory, the anisotropic ray theory, and the prevailing-frequency approximation of the coupling ray theory for S waves (Klimeš & Bulant, 2012).

## 2. Approximation of the complex-valued Green-tensor amplitude

Complex-valued Green-tensor amplitude  $\mathbf{G}^C$  maps seismic force  $\mathbf{f}$  onto the complex-valued displacement  $\mathbf{u}^C$ ,

$$\mathbf{u}^C = \mathbf{G}^C \mathbf{f} \quad . \quad (1)$$

We wish to approximate complex-valued Green-tensor amplitude  $\mathbf{G}^C$  by real-valued Green-tensor amplitude  $\mathbf{G}^R$  with phase shift factor  $\exp(i\varphi)$ , which map seismic force  $\mathbf{f}$  onto the approximate complex-valued displacement  $\mathbf{u}^R$ ,

$$\mathbf{u}^R = \mathbf{G}^R \mathbf{f} \exp(i\varphi) \quad . \quad (2)$$

The square of the wavefield error of this approximation is

$$\delta^2 = (\mathbf{u}^C - \mathbf{u}^R) + (\mathbf{u}^C - \mathbf{u}^R) \quad , \quad (3)$$

where  $^+$  denotes the Hermitian adjoint. We insert relations (1) and (2) into definition (3) and obtain

$$\delta^2 = \mathbf{f}^+ [\mathbf{G}^C - \mathbf{G}^R \exp(i\varphi)]^+ [\mathbf{G}^C - \mathbf{G}^R \exp(i\varphi)] \mathbf{f} \quad . \quad (4)$$

We average the square of wavefield error  $\delta^2$  over all spatial directions of unit seismic force  $\mathbf{f}$ ,

$$\langle \delta^2 \rangle = \frac{1}{3} \text{Tr}\{ [\mathbf{G}^C - \mathbf{G}^R \exp(i\varphi)]^+ [\mathbf{G}^C - \mathbf{G}^R \exp(i\varphi)] \} \quad . \quad (5)$$

After multiplication, relation (5) reads

$$\langle \delta^2 \rangle = \frac{1}{3} \text{Tr}[(\mathbf{G}^C)^+ \mathbf{G}^C - (\mathbf{G}^R)^T \mathbf{G}^C \exp(-i\varphi) - (\mathbf{G}^C)^+ \mathbf{G}^R \exp(i\varphi) + (\mathbf{G}^R)^T \mathbf{G}^R] \quad , \quad (6)$$

where  $^T$  denotes the transpose. We decompose the complex-valued Green-tensor amplitude into its real and imaginary parts and obtain

$$\langle \delta^2 \rangle = \frac{1}{3} \text{Tr}[(\mathbf{G}^C)^+ \mathbf{G}^C - 2(\mathbf{G}^R)^T \text{Re}(\mathbf{G}^C) \cos(\varphi) - 2(\mathbf{G}^R)^T \text{Im}(\mathbf{G}^C) \sin(\varphi) + (\mathbf{G}^R)^T \mathbf{G}^R] \quad . \quad (7)$$

### 3. Minimizing the wavefield error

We wish to select real-valued Green-tensor amplitude  $\mathbf{G}^R$  and phase shift  $\varphi$  so that the average square (7) of wavefield error is minimal.

Differentiating wavefield error (7) with respect to real-valued Green-tensor amplitude  $\mathbf{G}^R$ , we obtain equation

$$\frac{2}{3} [-\text{Re}(\mathbf{G}^C) \cos(\varphi) - \text{Im}(\mathbf{G}^C) \sin(\varphi) + \mathbf{G}^R] = \mathbf{0} \quad (8)$$

for the minimum of the average square (7) of wavefield error.

Differentiating wavefield error (7) with respect to phase shift  $\varphi$ , we obtain equation

$$\frac{2}{3} \text{Tr}[(\mathbf{G}^R)^T \text{Re}(\mathbf{G}^C) \sin(\varphi) - (\mathbf{G}^R)^T \text{Im}(\mathbf{G}^C) \cos(\varphi)] = 0 \quad . \quad (9)$$

for the minimum of the average square (7) of wavefield error.

Equation (8) yields

$$\mathbf{G}^R = \text{Re}(\mathbf{G}^C) \cos(\varphi) + \text{Im}(\mathbf{G}^C) \sin(\varphi) \quad . \quad (10)$$

We insert relation (10) into equation (9) and arrive at

$$\begin{aligned} \frac{2}{3} \text{Tr}[\text{Re}(\mathbf{G}^C)^T \text{Re}(\mathbf{G}^C) \cos(\varphi) \sin(\varphi) + \text{Im}(\mathbf{G}^C)^T \text{Re}(\mathbf{G}^C) \sin(\varphi) \sin(\varphi) \\ - \text{Re}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C) \cos(\varphi) \cos(\varphi) - \text{Im}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C) \sin(\varphi) \cos(\varphi)] = 0 \quad , \quad (11) \end{aligned}$$

which yields

$$\text{Tr}[\text{Re}(\mathbf{G}^C)^T \text{Re}(\mathbf{G}^C) - \text{Im}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C)] \sin(2\varphi) = 2 \text{Tr}[\text{Re}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C)] \cos(2\varphi) \quad . \quad (12)$$

The solution of equation (12) is

$$\tan(2\varphi) = 2 \text{Tr}[\text{Re}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C)] / \text{Tr}[\text{Re}(\mathbf{G}^C)^T \text{Re}(\mathbf{G}^C) - \text{Im}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C)] \quad . \quad (13)$$

We can thus calculate phase shift  $\varphi$  using relation (13) and insert it into relation (10) to obtain the best real-valued Green-tensor amplitude  $\mathbf{G}^R$ .

#### 4. Relative wavefield error

For the best real-valued Green-tensor amplitude (10), the average square (7) of wavefield error reads

$$\langle \delta^2 \rangle = \frac{1}{3} \text{Tr}[(\mathbf{G}^C)^+ \mathbf{G}^C - (\mathbf{G}^R)^T \mathbf{G}^R] \quad . \quad (14)$$

The square of the norm of the complex-valued displacement is

$$|\mathbf{u}^C|^2 = \mathbf{f}^+ (\mathbf{G}^C)^+ \mathbf{G}^C \mathbf{f} \quad . \quad (15)$$

We average the square of the norm over all spatial directions of unit seismic force  $\mathbf{f}$ ,

$$\langle |\mathbf{u}^C|^2 \rangle = \frac{1}{3} \text{Tr}[(\mathbf{G}^C)^+ \mathbf{G}^C] \quad . \quad (16)$$

The relative root mean square error of the approximation of the complex-valued Green-tensor amplitude by the real-valued Green-tensor amplitude is

$$\rho = \sqrt{\langle \delta^2 \rangle / \langle |\mathbf{u}^C|^2 \rangle} \quad , \quad (17)$$

where  $\langle \delta^2 \rangle$  and  $\langle |\mathbf{u}^C|^2 \rangle$  are given by expressions (14) and (16) with

$$\text{Tr}[(\mathbf{G}^C)^+ \mathbf{G}^C] = \text{Tr}[\text{Re}(\mathbf{G}^C)^T \text{Re}(\mathbf{G}^C)] + \text{Tr}[\text{Im}(\mathbf{G}^C)^T \text{Im}(\mathbf{G}^C)] \quad . \quad (18)$$

If relative error (17) is acceptable, the real-valued Green-tensor amplitude can be used in the inversion according to Šílený & Milev (2008), Šílený et al. (2009), and Horálek & Šílený (2013). If relative error (17) is not acceptable, the corresponding data from the receiver should not be used in that inversion.

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#### References

- Červený, V. (2001): *Seismic Ray Theory*. Cambridge Univ. Press, Cambridge.
- Horálek, J. & Šílený, J. (2013): Source mechanisms of the 2000 earthquake swarm in the West Bohemia/Vogtland region (Central Europe). *Geophys. J. Int.*, **194**, 979-999.
- Klimeš, L. & Bulant, P. (2012): Single-frequency approximation of the coupling ray theory. In: *Seismic Waves in Complex 3-D Structures, Report 22*, pp. 143-167, Dep. Geophys., Charles Univ., Prague, online at “<http://sw3d.cz>”.
- Šílený, J., Hill, D.P., Eisner, L. & Cornet F.H. (2009): Non-double-couple mechanisms of microearthquakes induced by hydraulic fracturing. *J. Geophys. Res.*, **114B**, 08307.
- Šílený, J. & Milev, A. (2008): Source mechanism of mining induced seismic events — resolution of double couple and non double couple models. *Tectonophysics*, **456**, 3-15.